

Seasonal Variability of Wave States and Their Influence on Air-Sea Fluxes in the Northern Indian Ocean

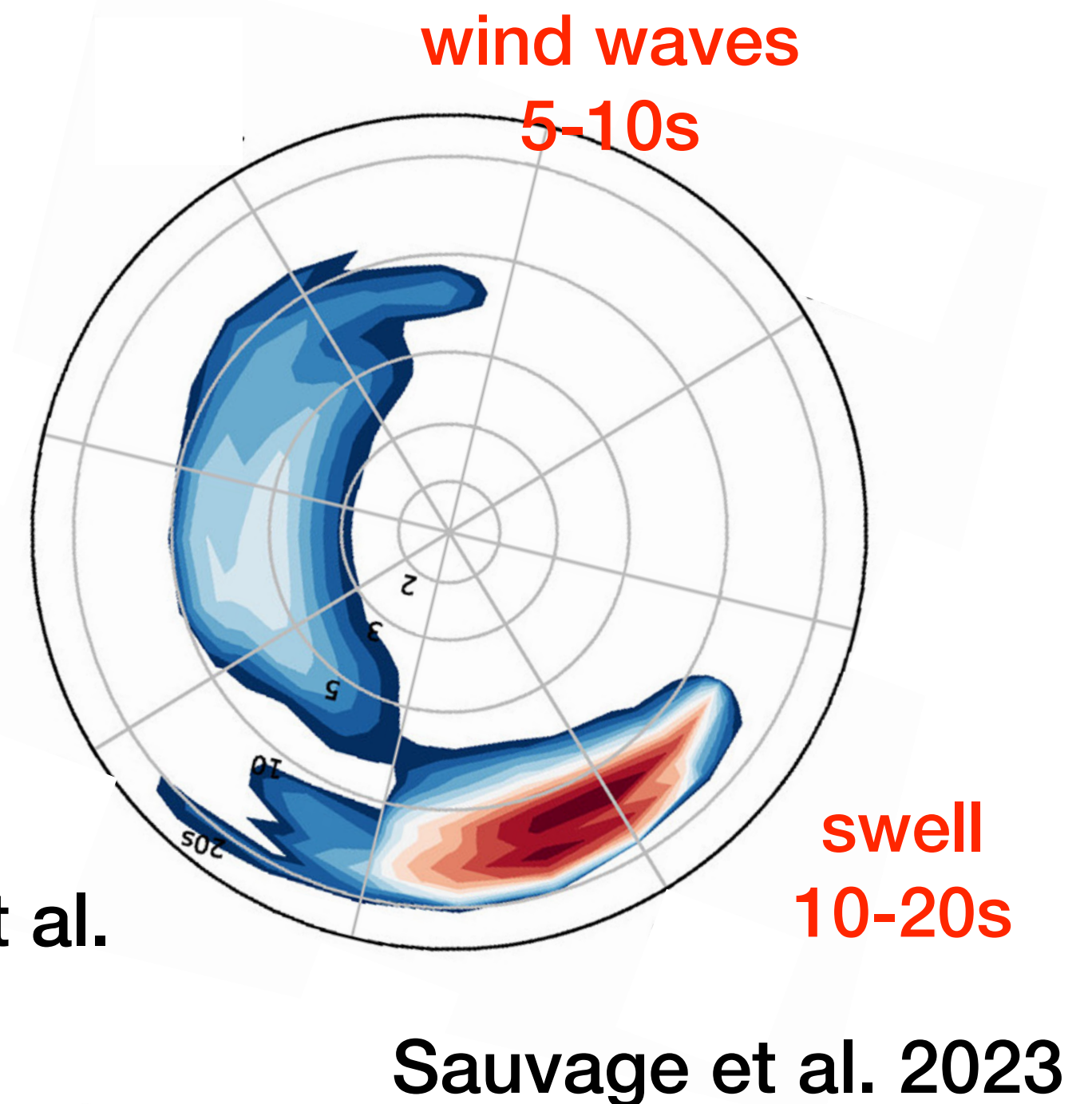
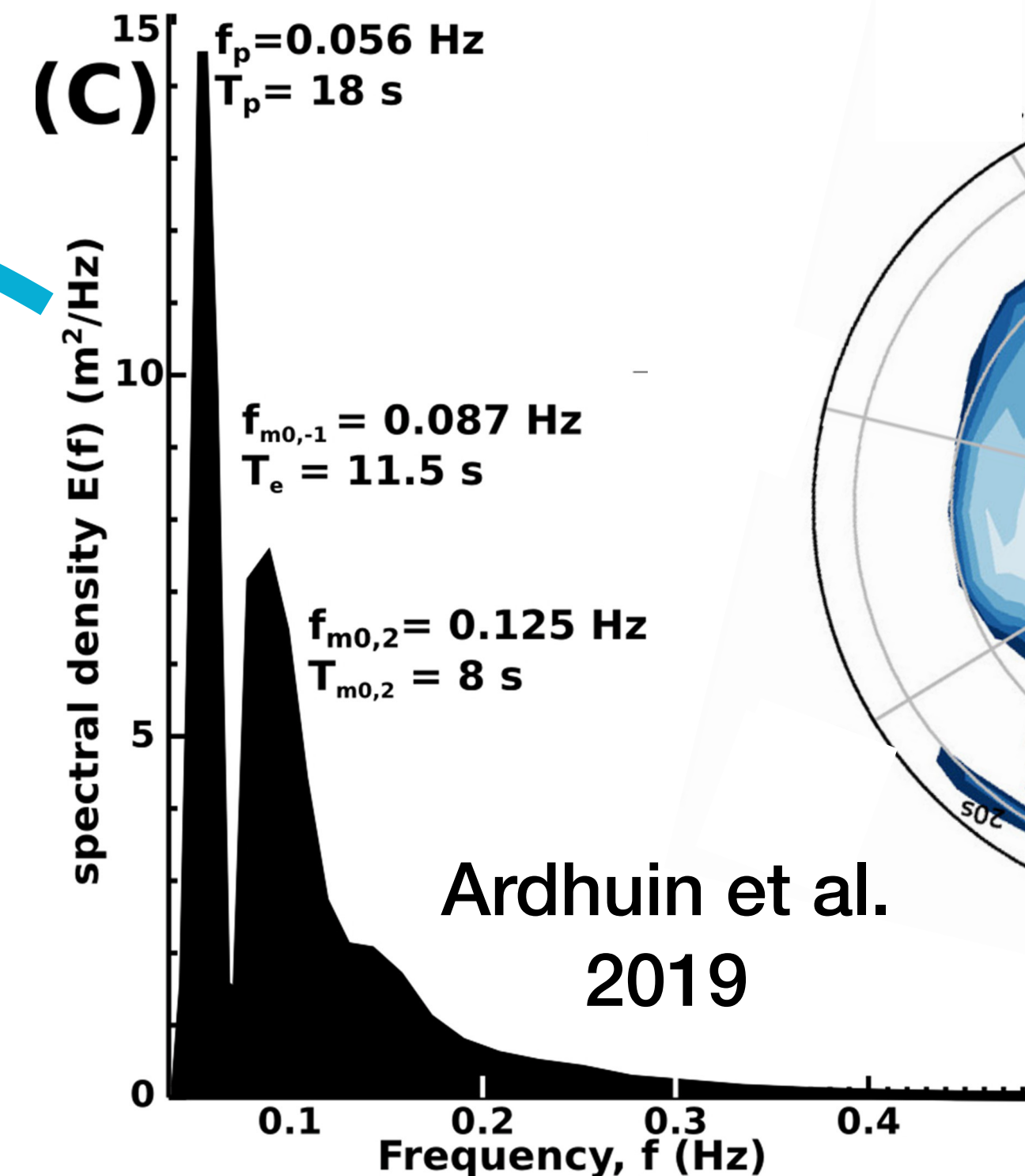
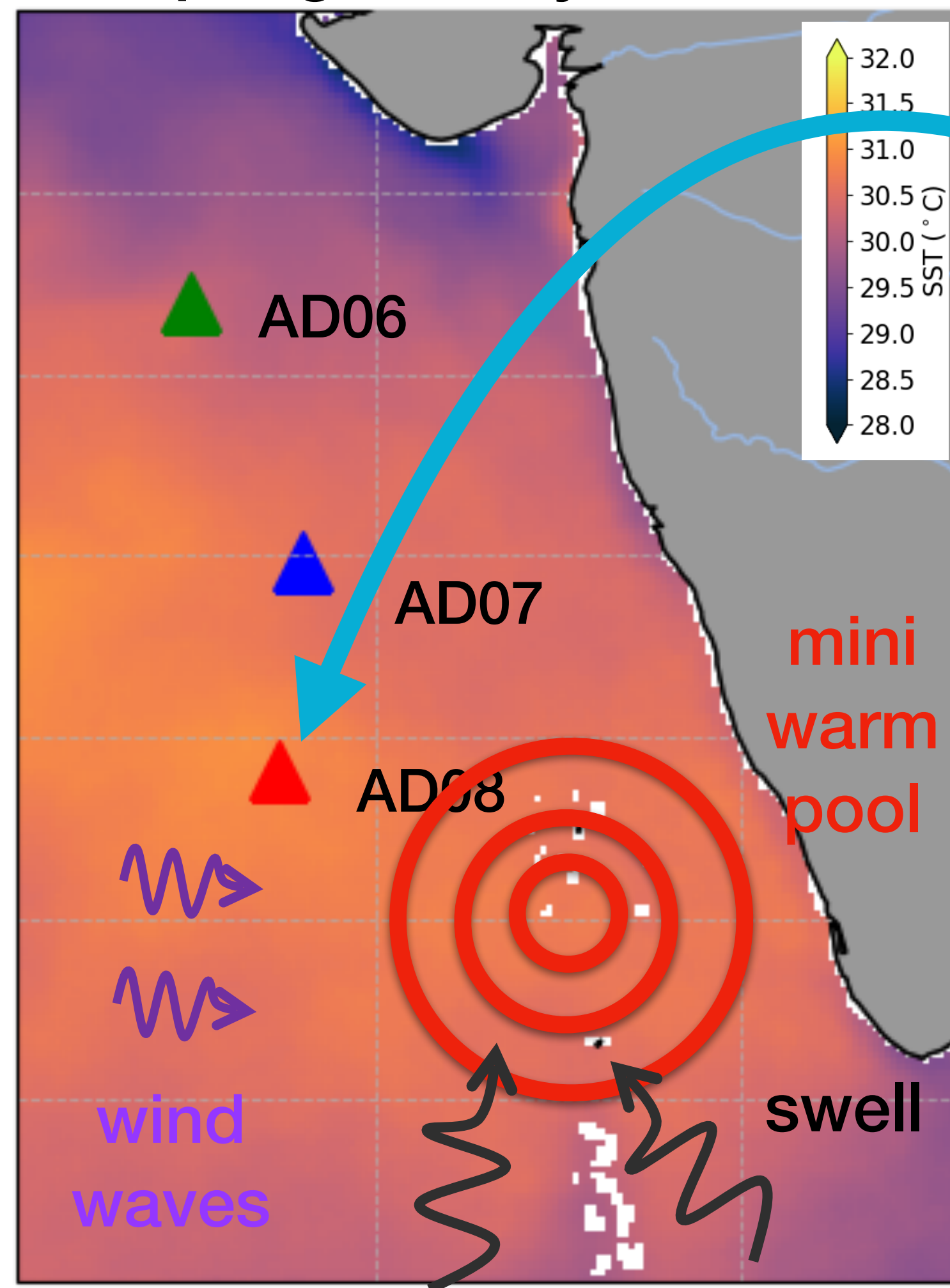


OMNI mooring



Venkatesan et al. 2021

Spring to early summer

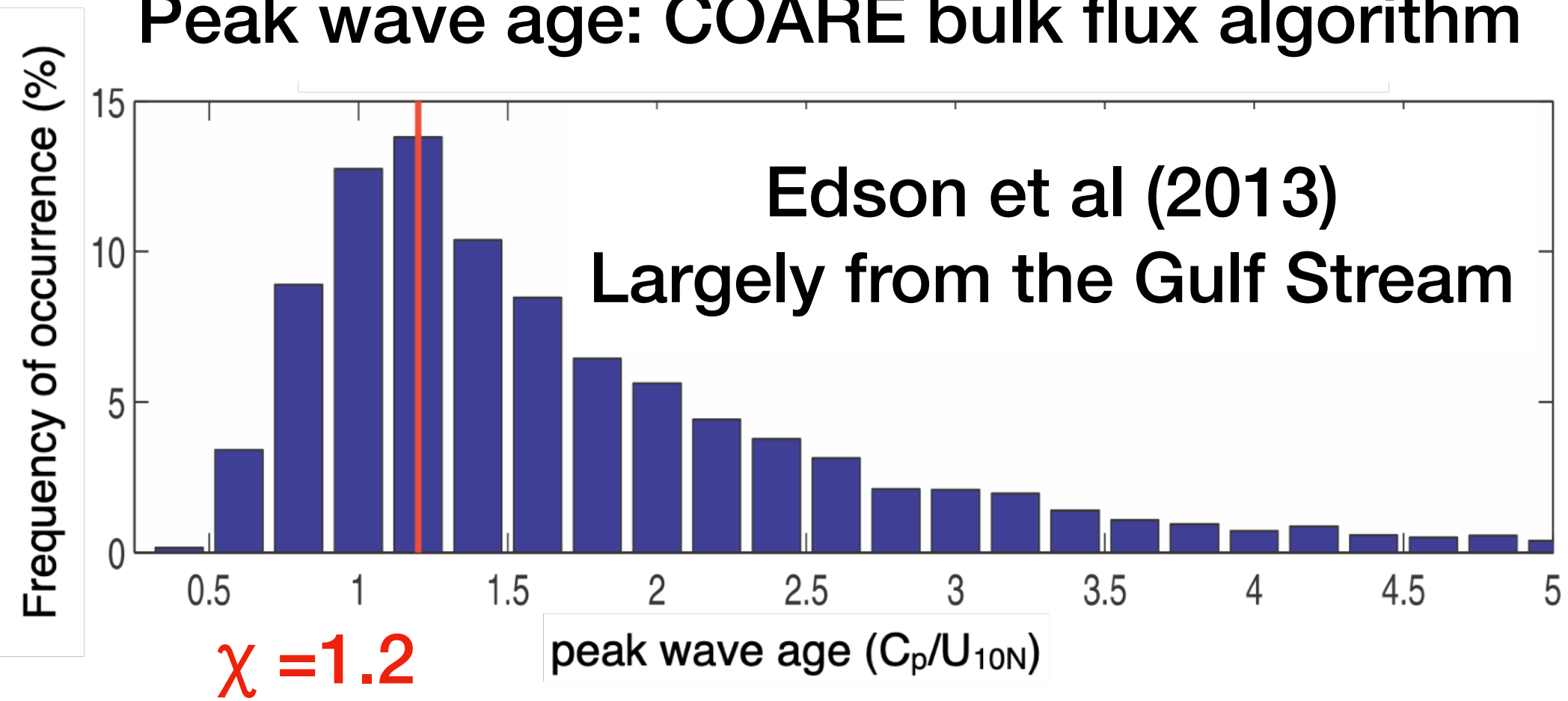


Hyodae Seo, Sid Kerhalkar, César Sauvage (Univ. Hawaii, Manoa), Amit Tandon (UMass Dartmouth)

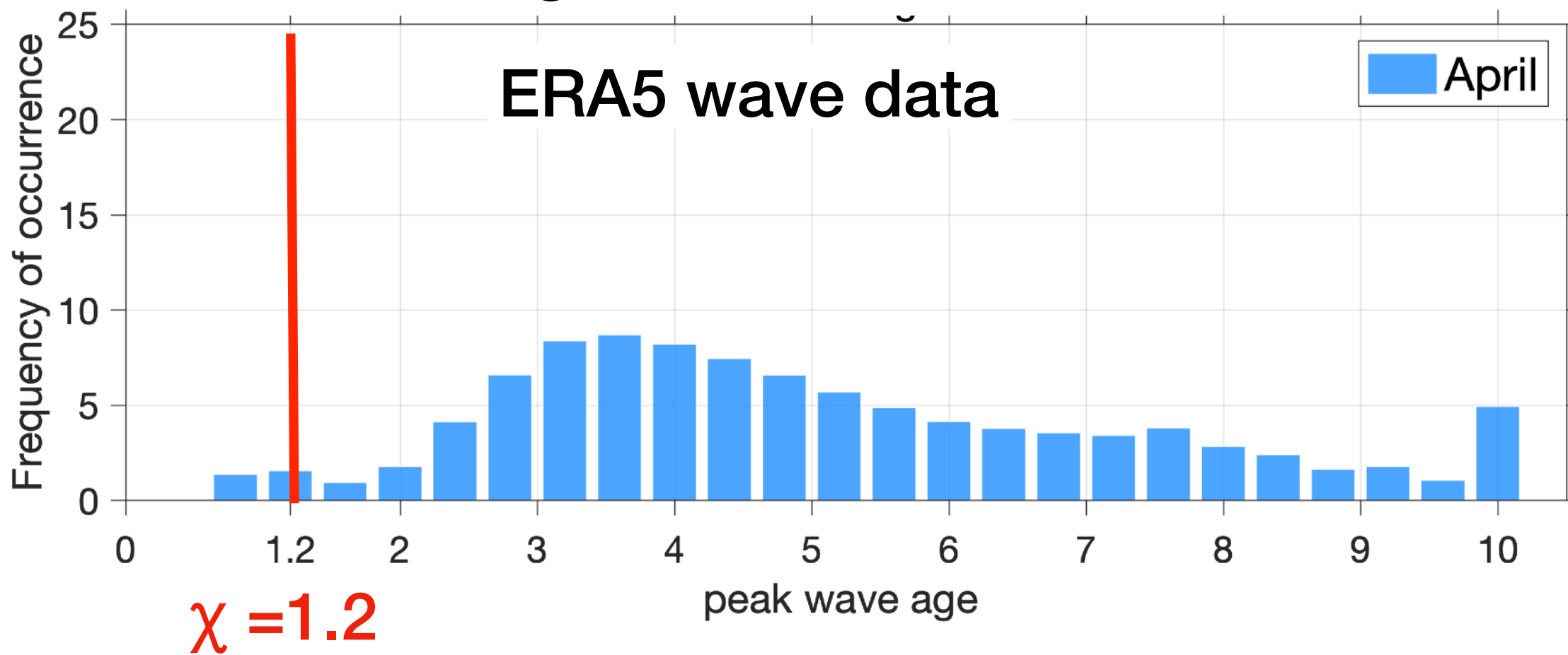
(Peak) Wave age

$$\chi \equiv C_p/U_{10}$$

Peak wave age: COARE bulk flux algorithm

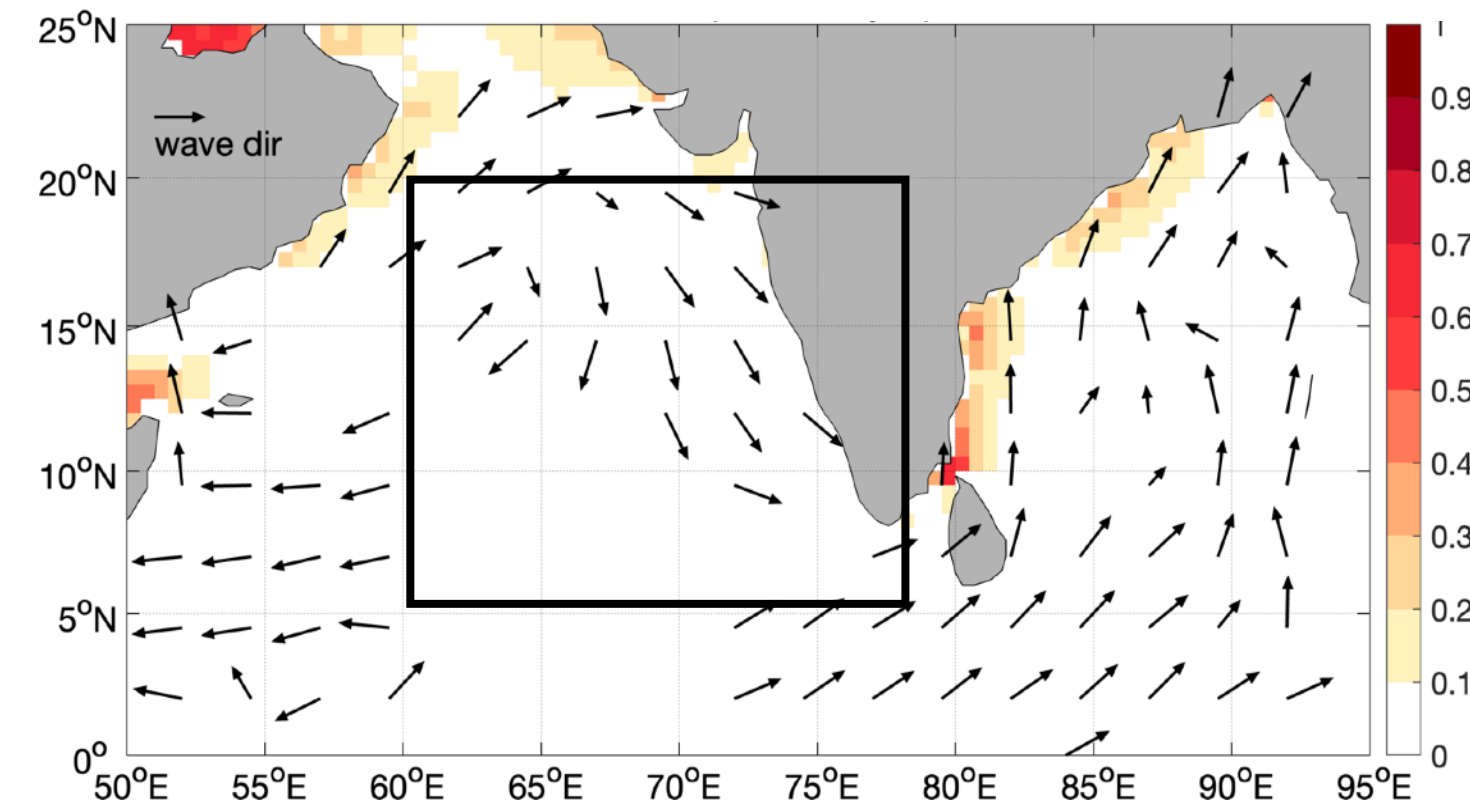


Peak wave age: Arabian Sea

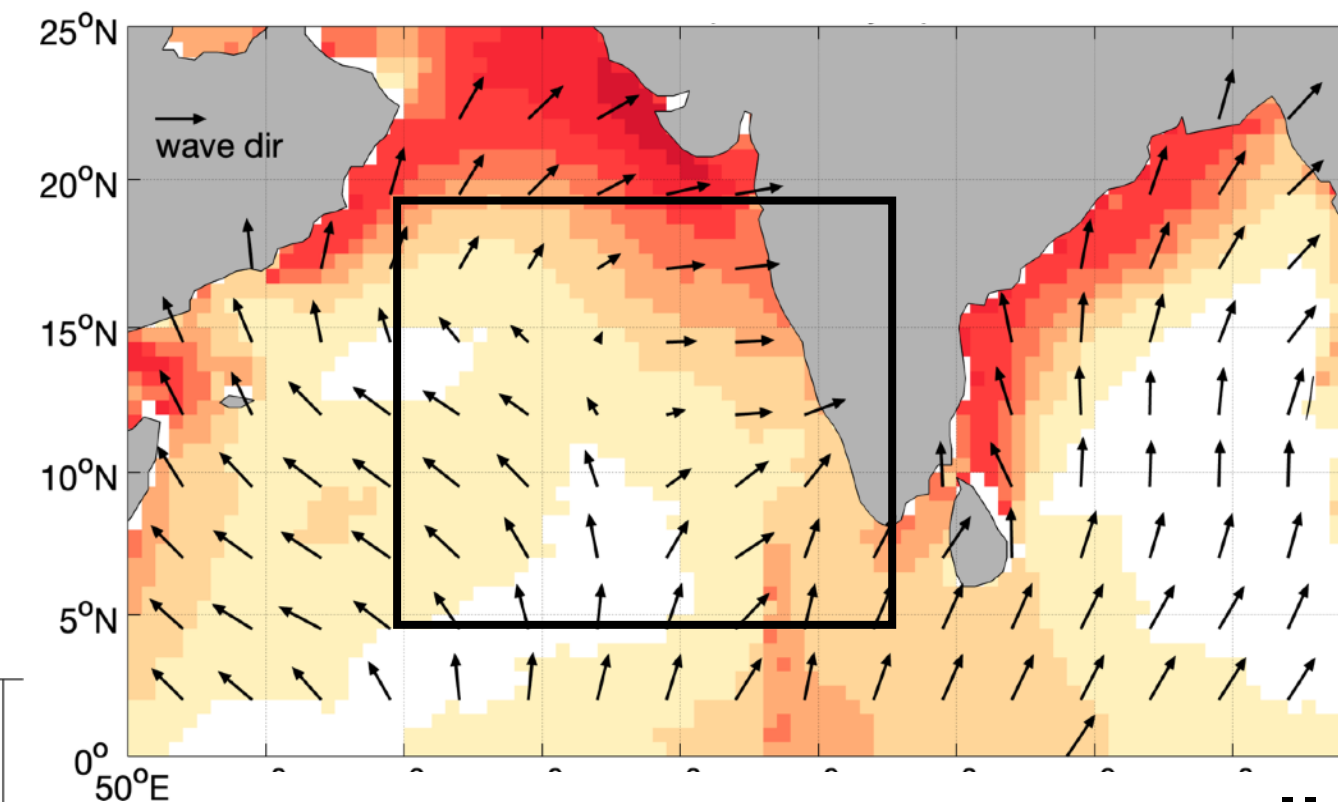


April

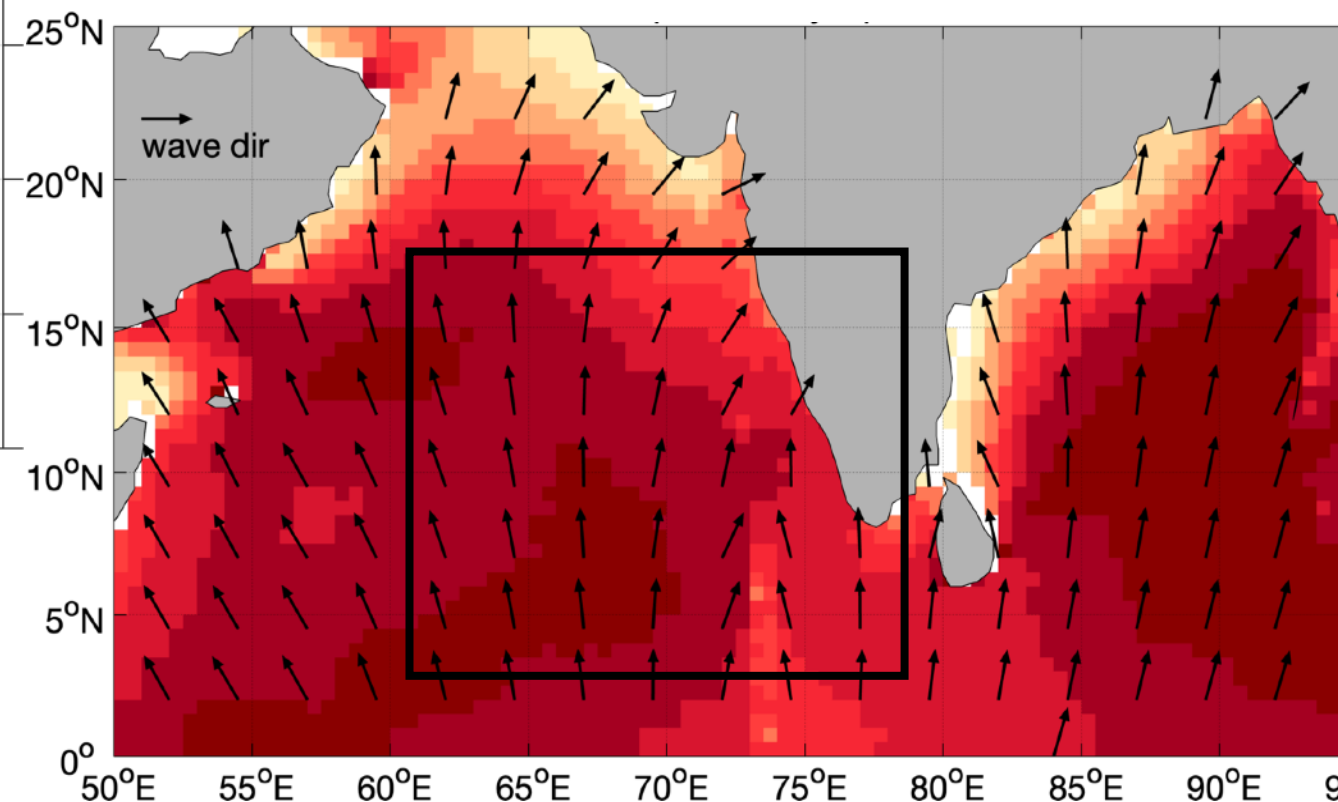
windsea ($\chi < 1.2$)



mixed sea ($1.2 < \chi < 3$)



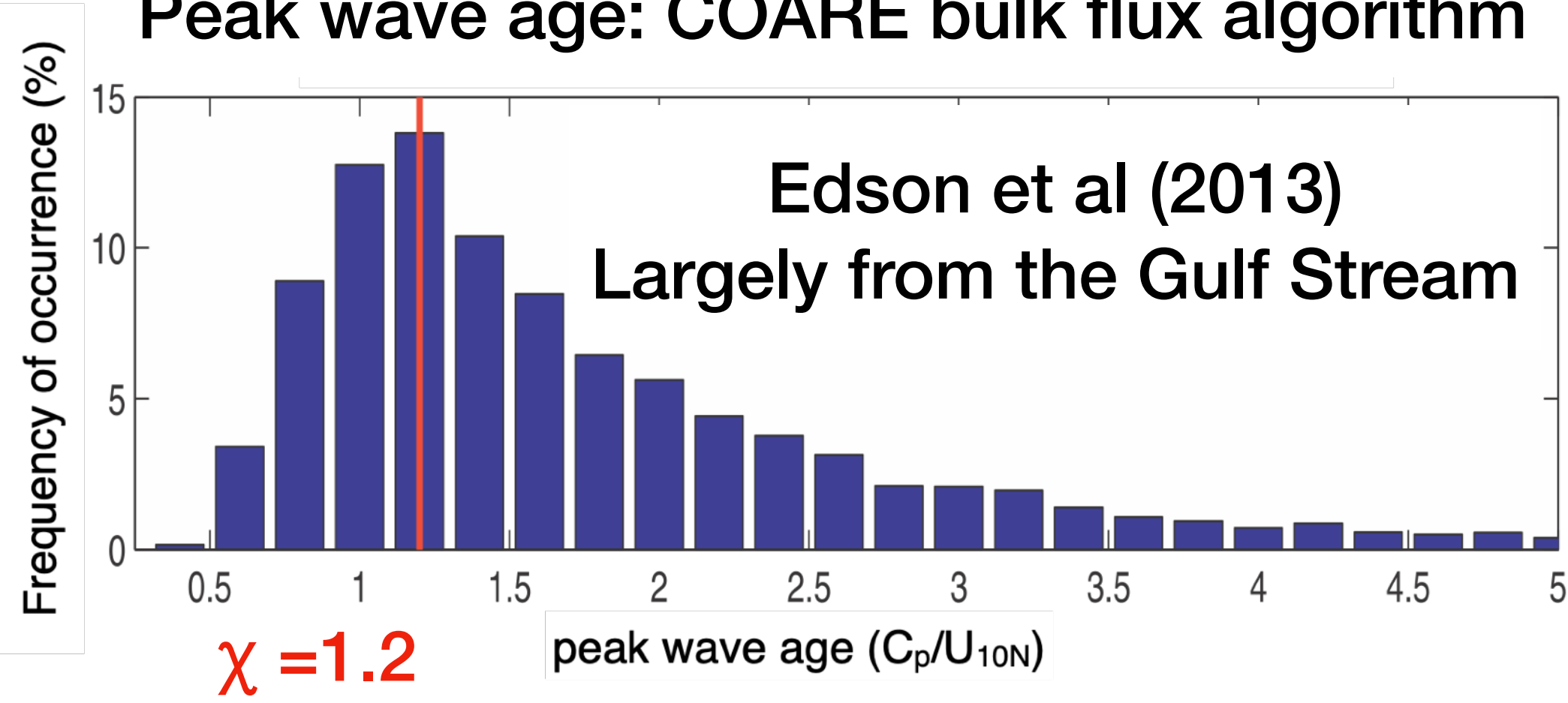
swell ($\chi > 3$)



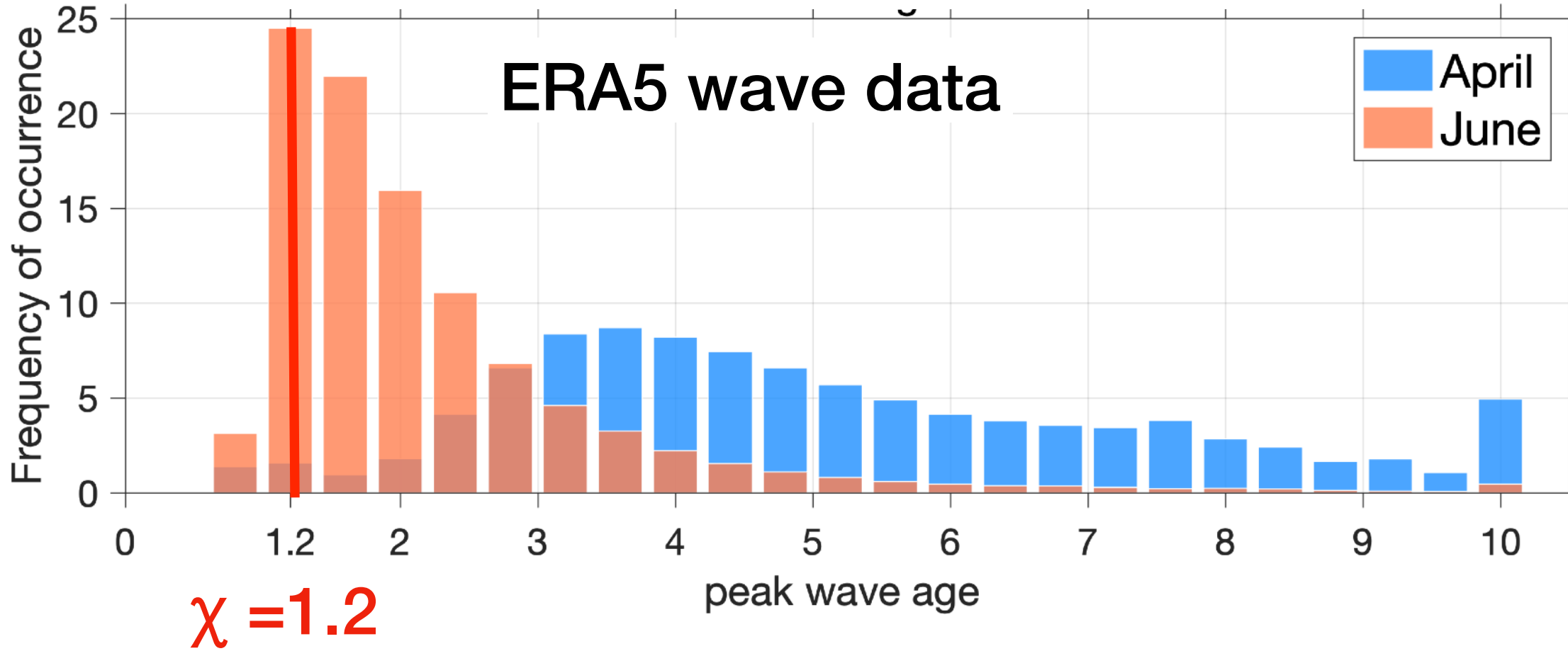
(Peak) Wave age

$\chi \equiv C_p/U_{10}$ (or C_p/u^*)

Peak wave age: COARE bulk flux algorithm



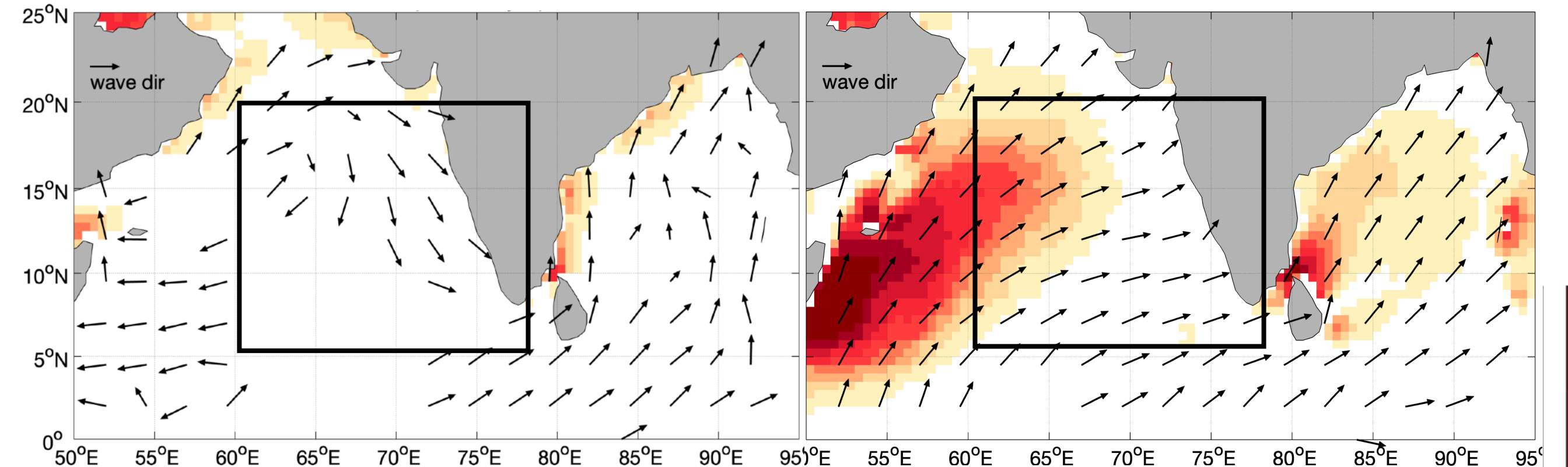
Peak wave age: Arabian Sea



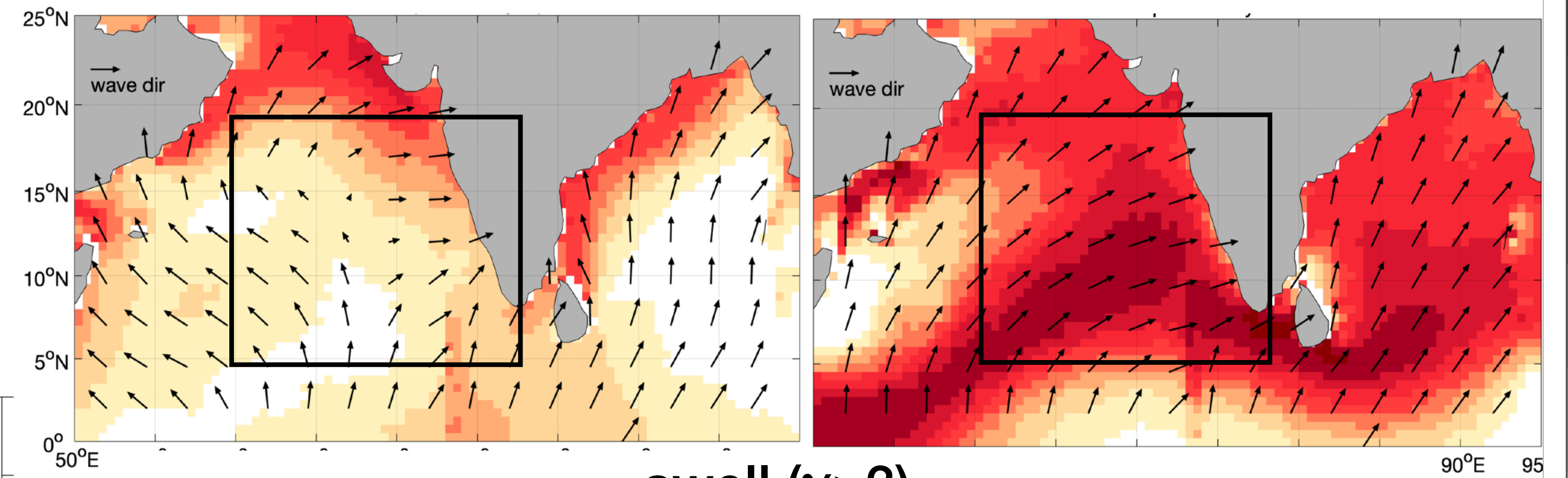
April

windsea ($\chi < 1.2$)

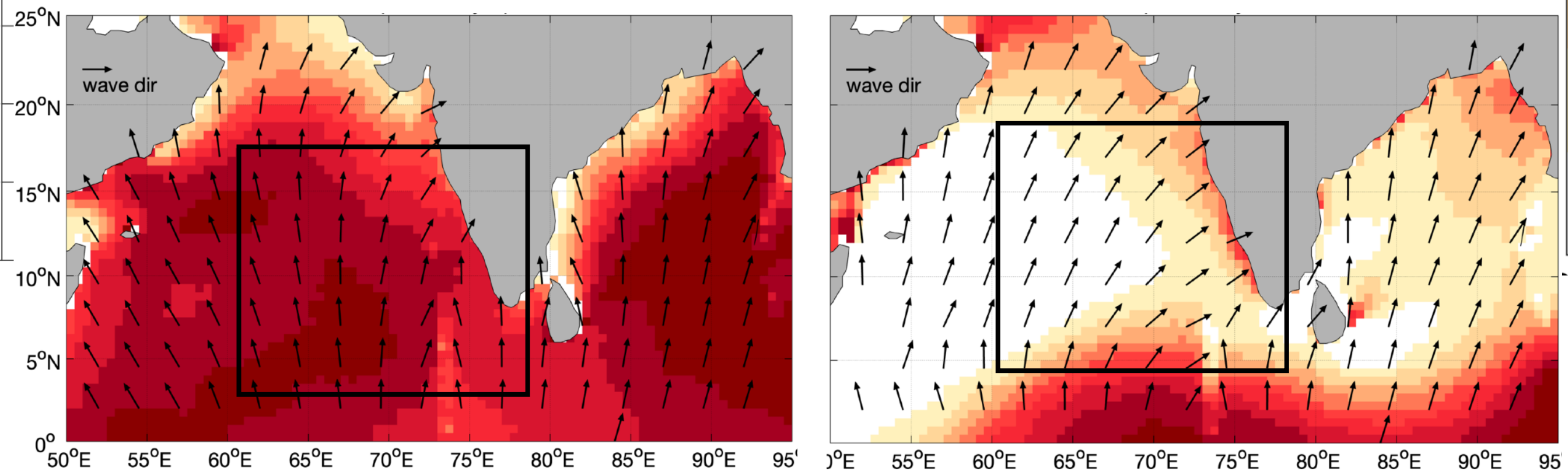
June



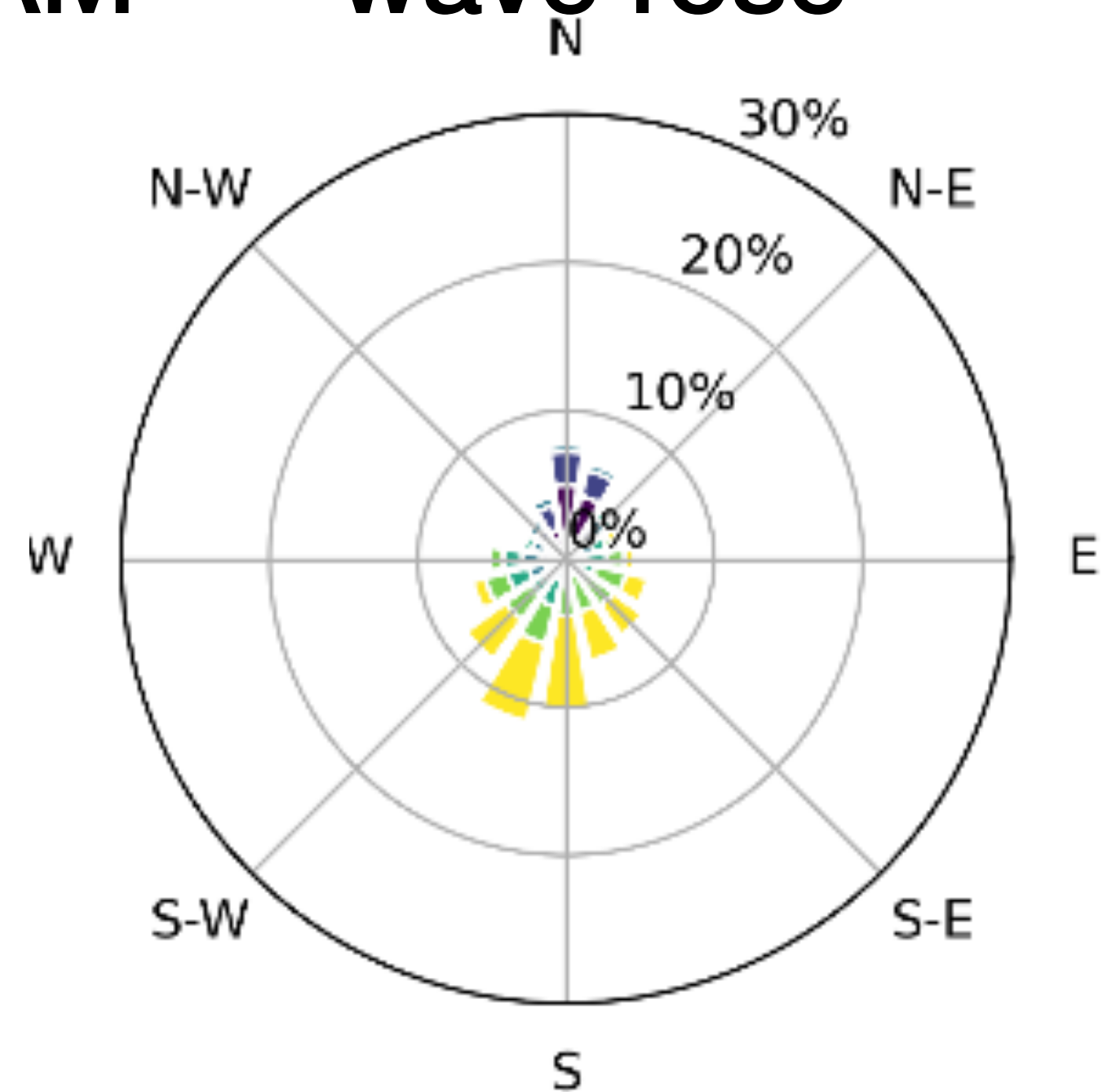
mixed sea ($1.2 < \chi < 3$)



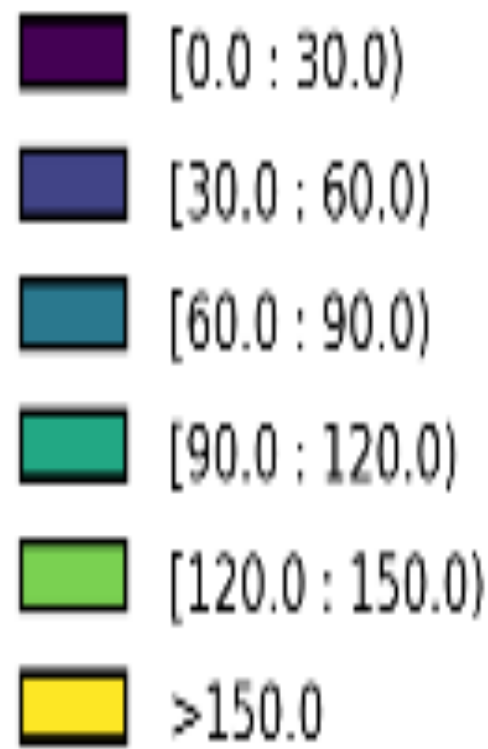
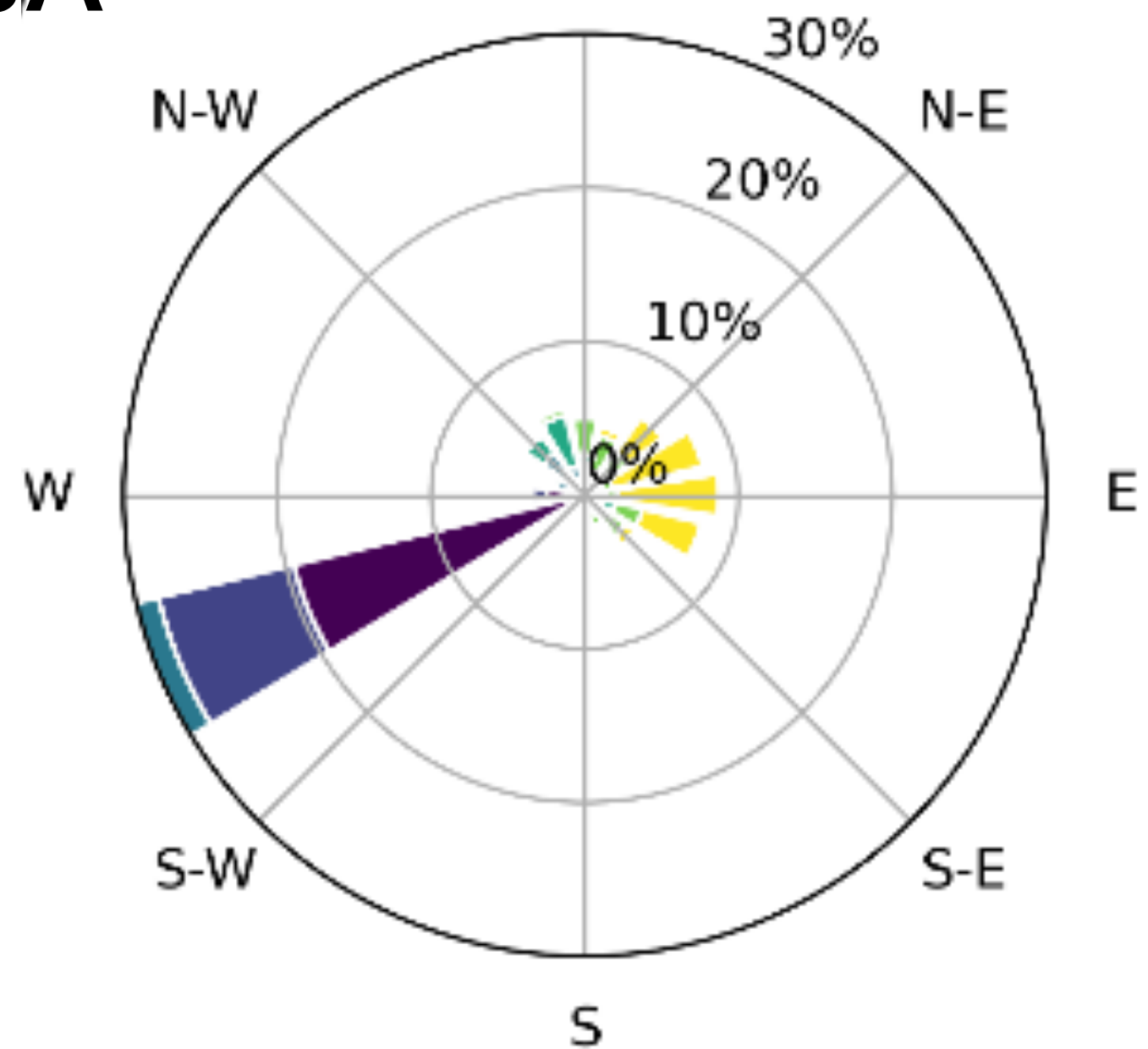
swell ($\chi > 3$)



MAM wave rose

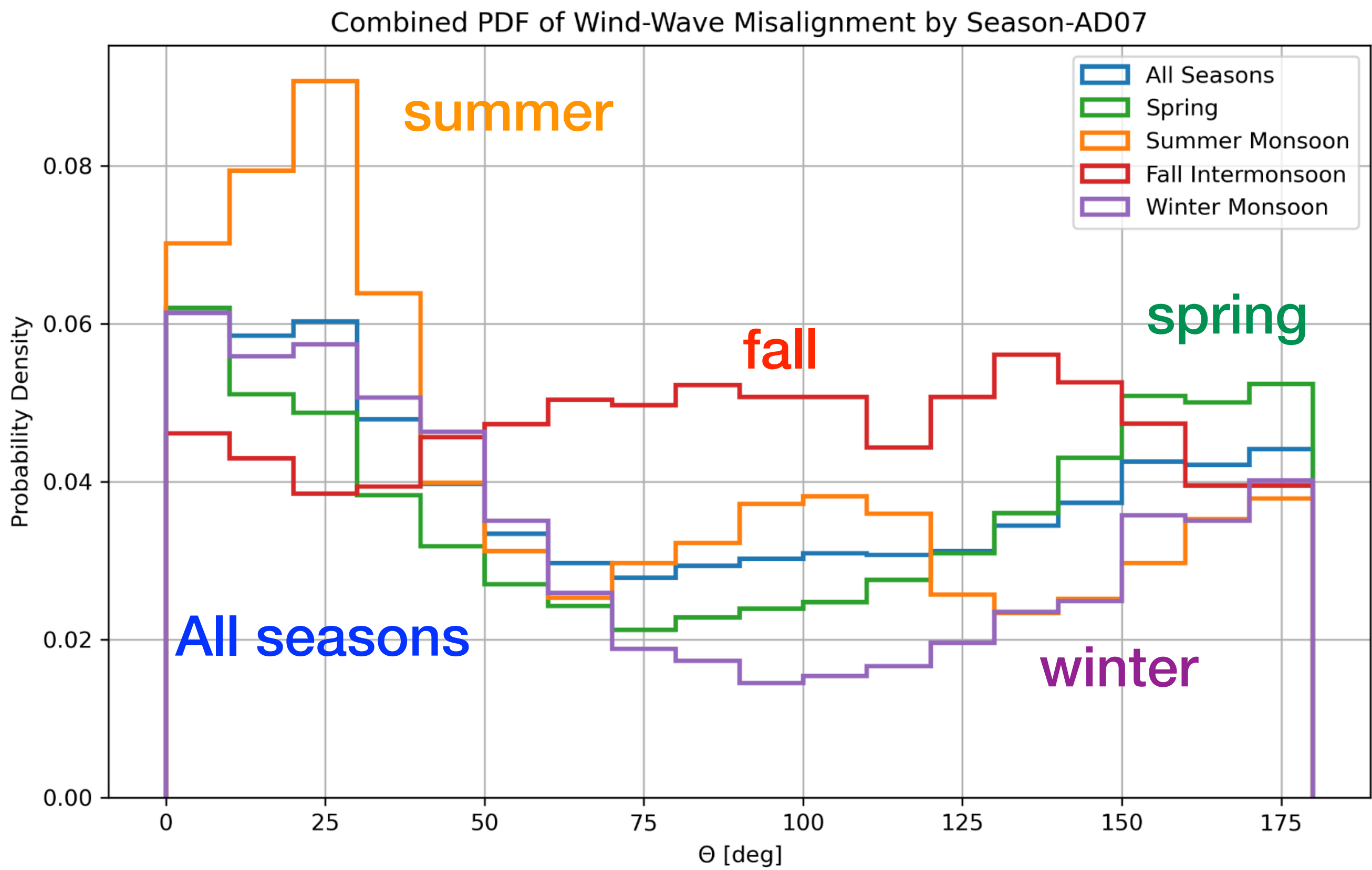


JJA



Observed wave-wind misalignment at AD07

See also Raj et al. (2025)

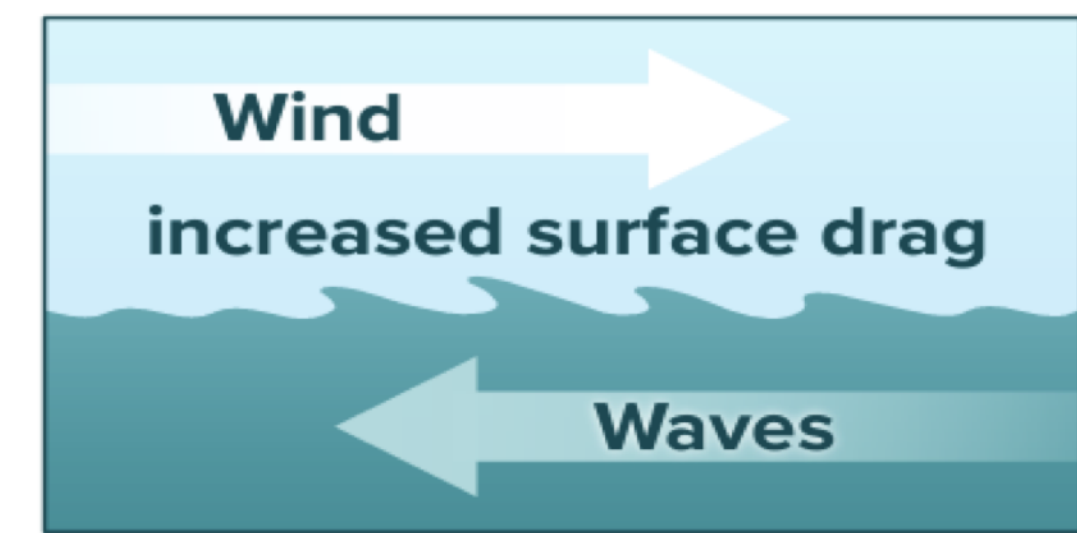


Misaligned waves observed ubiquitously throughout the year

OMNI mooring AD07 Hourly wind and wave data: 2012-2024

How is surface roughness parameterized?

COARE algorithm (Edson et al. 2013)



$$\tau = \rho_a \underbrace{C_D}_{\text{roughness}} U_r^2 \quad C_D \cong \left[\frac{\kappa}{\ln(z/z_0) - \psi_m(z/L)} \right]^2 \quad z_0 = z_0^{\text{smooth}} + z_0^{\text{rough}}$$

1. Wind Speed Dependent Formulation (WSDF)

$$z_0^{\text{rough}} = \alpha \frac{u_*^2}{g} \quad \alpha = m U_{10N} + b$$

Charnock coefficient

- Assumes wind-wave equilibrium ($\chi \sim 1.2$)
- Assumes waves are aligned with winds ($\theta=0$)**
- Violated in mixed seas, as well as near strong density/vorticity fronts, shallow fetch-limited oceans, and rapidly translating cyclones

2. Wave-Based Formulation (WBF)

$$z_0^{\text{rough}} = H_s \cdot 0.09 \cdot \left(\frac{u_*}{C_p} \right)^2$$

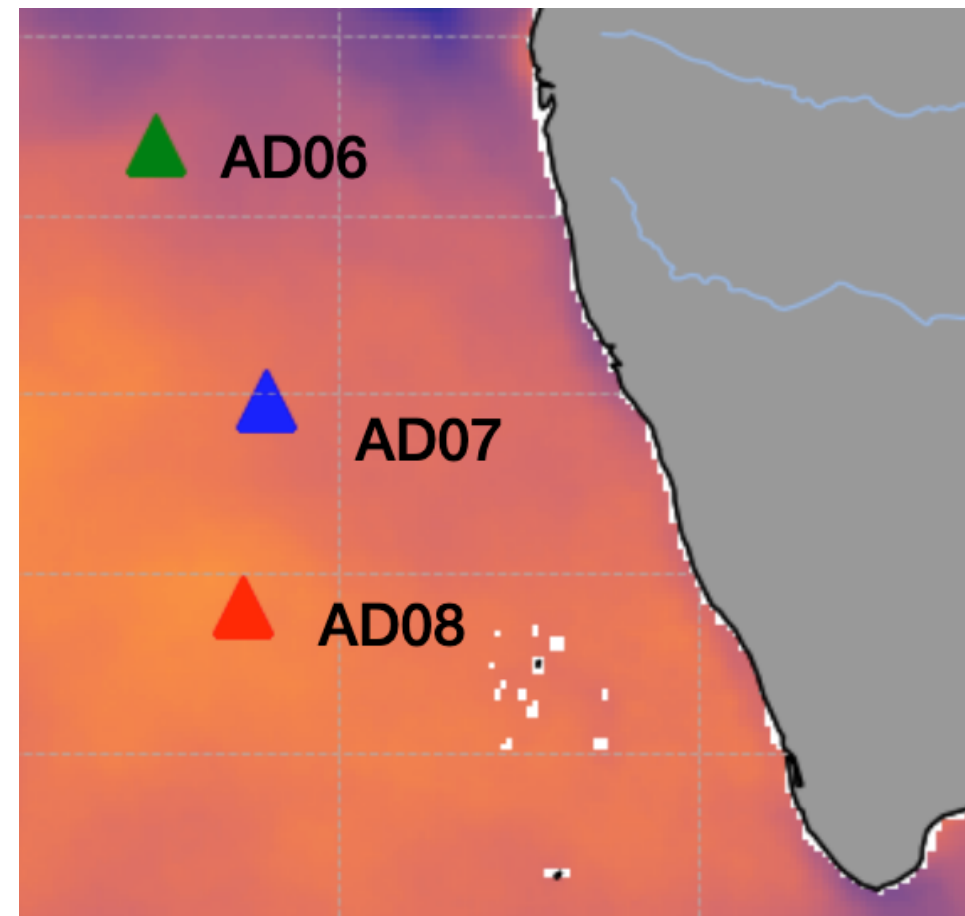
significant wave height

inverse
"peak" wave age

- Still assumes $\theta=0$** Sauvage et al. (2023, 2024)

$$\text{WBF}_{\theta} \quad z_0 = H_s \cdot 0.09 \cdot \cos(0.45 \theta) \cdot \left(\frac{u_*}{C_p} \right)^{2 \cdot \cos(-0.32 \theta)}$$

- Based on geographically biased wave datasets (e.g., no wave data from the low latitudes)

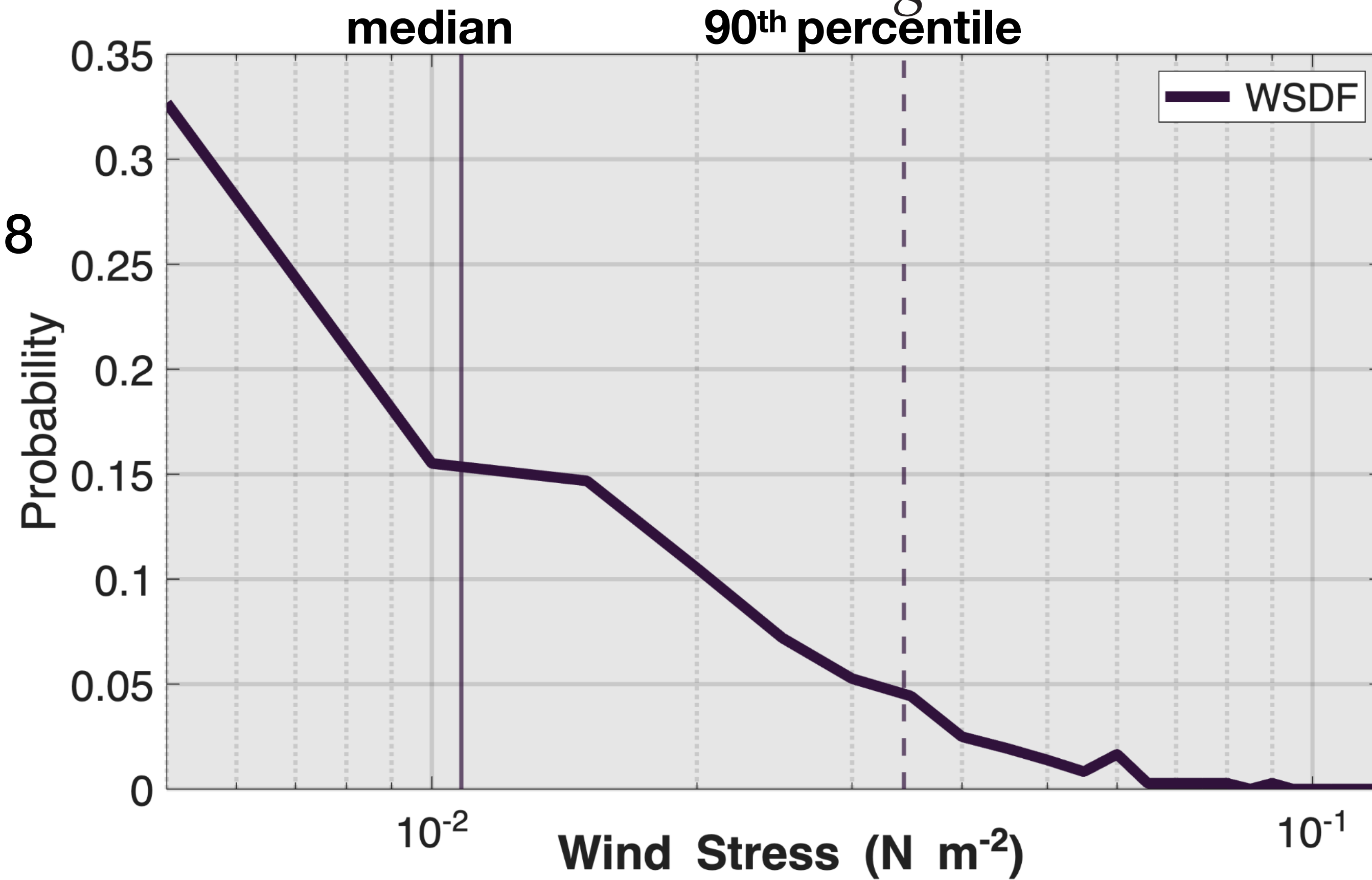


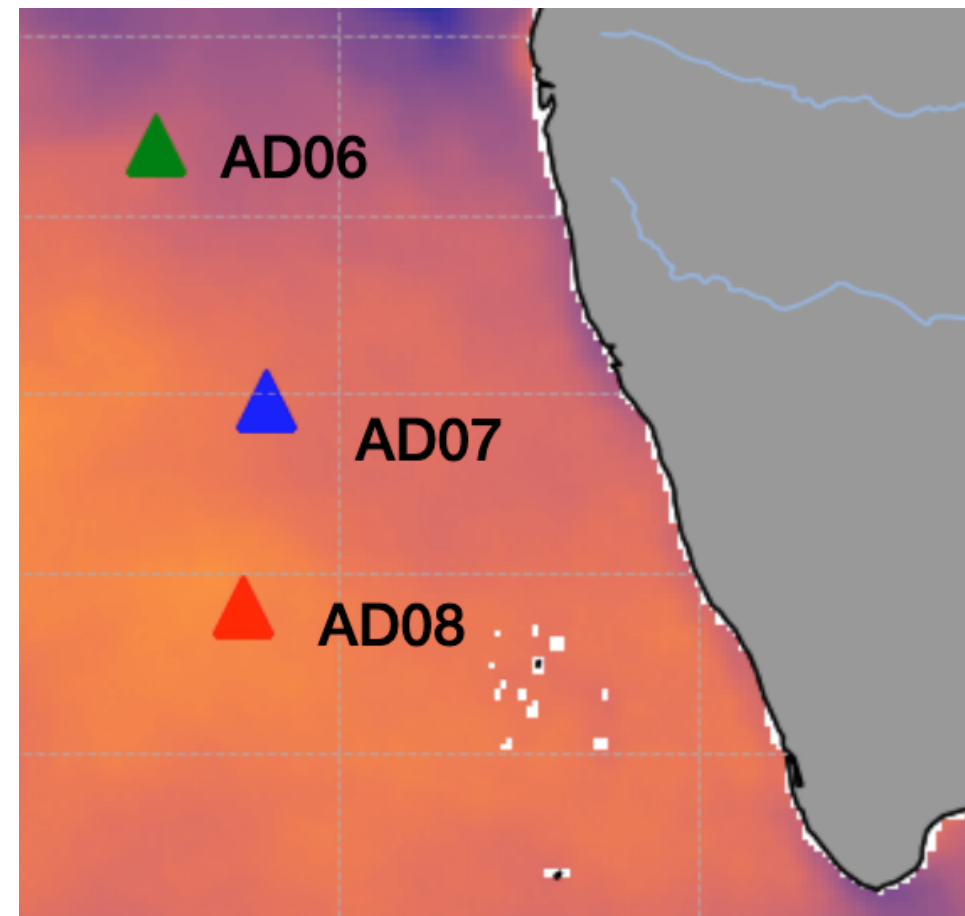
Wind stres estimated at AD08

$$z_0^{\text{rough}} = \alpha \frac{u_*^2}{g}$$

WSDF: Edson et al. (2013)

April 26— May 10, 2018



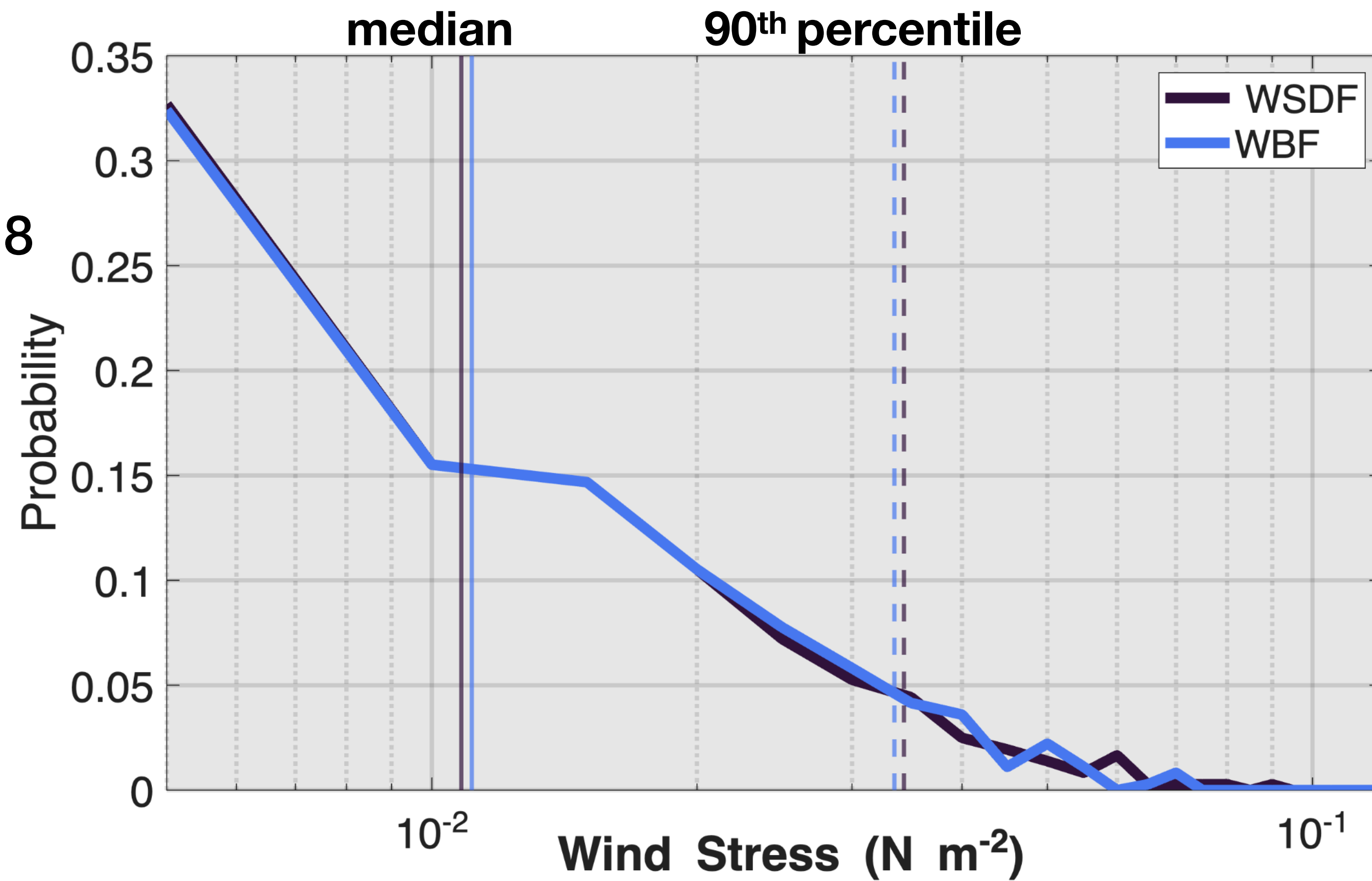


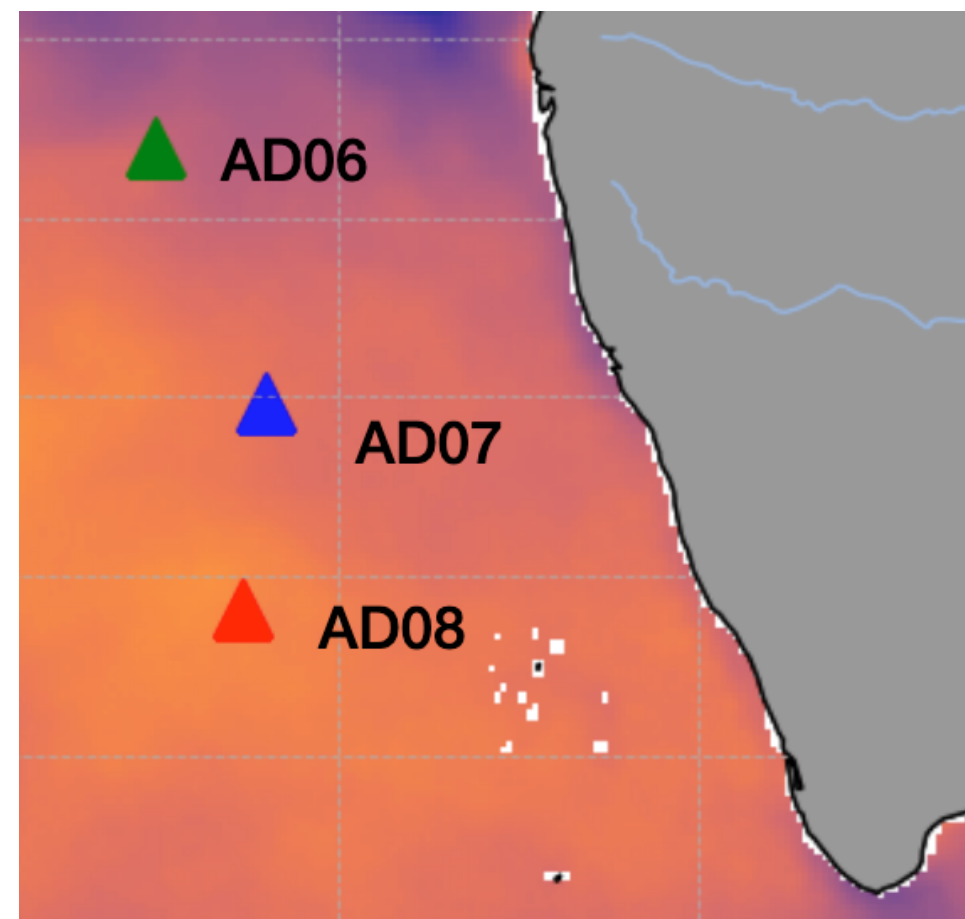
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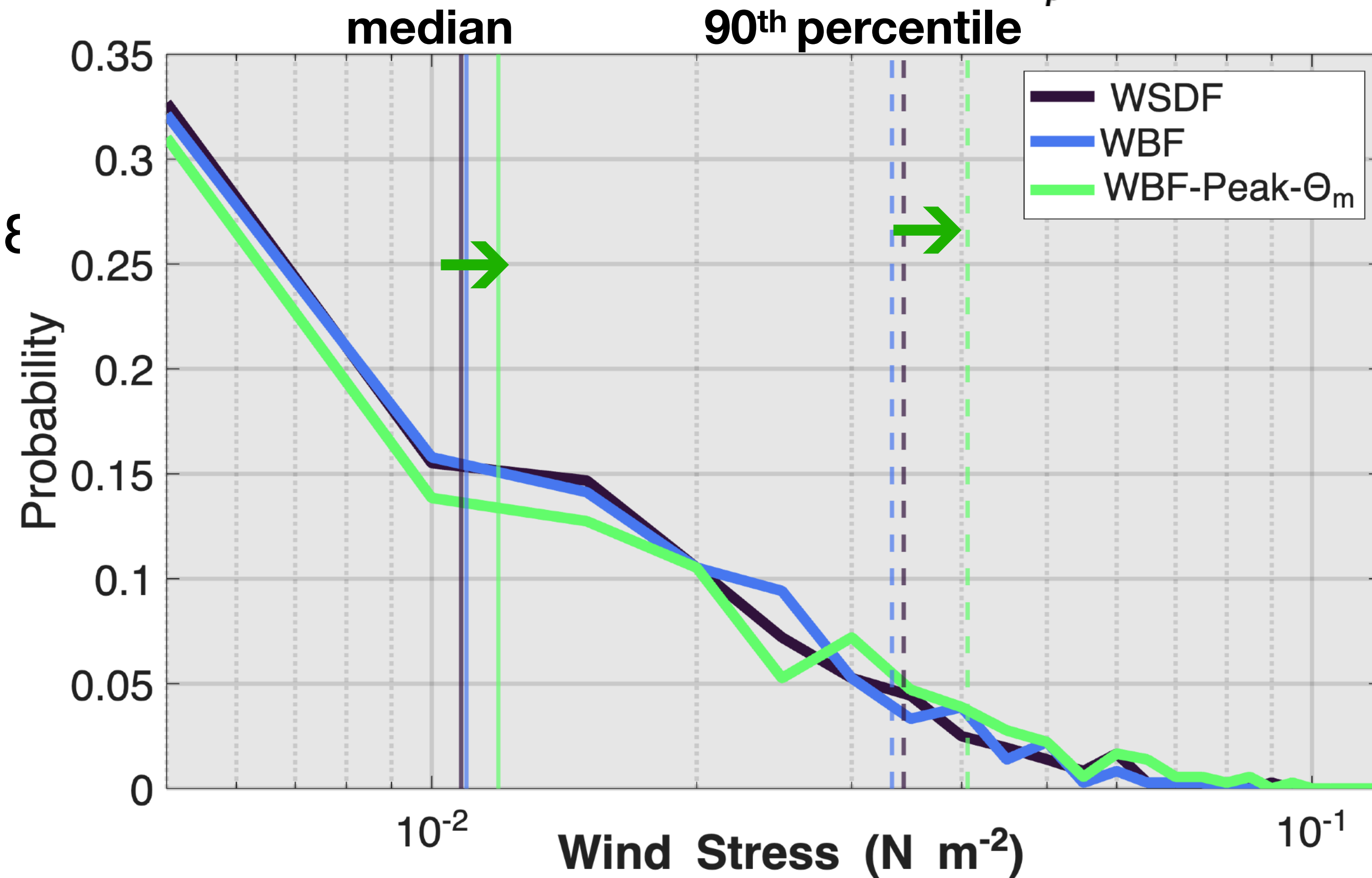


Wind stress estimated at AD08

$$Z_0^{rough} = H_s \cdot 0.09 \cdot \cos(0.45 \theta) \cdot \left(\frac{u_*}{C_p} \right)^{2 \cdot \cos(-0.32 \theta)}$$

WBF θ : Sauvage et al. (2023)

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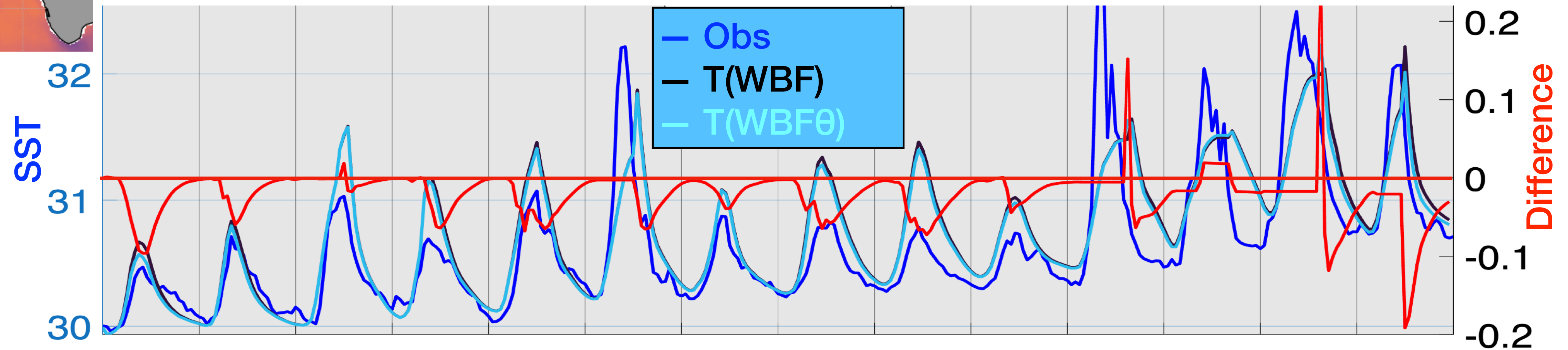
← Wind Stress is increased by 10-20% if accounting for wind-wave misalignment.

- There is no publically available direct covariance flux dataset to validate these parameterizations

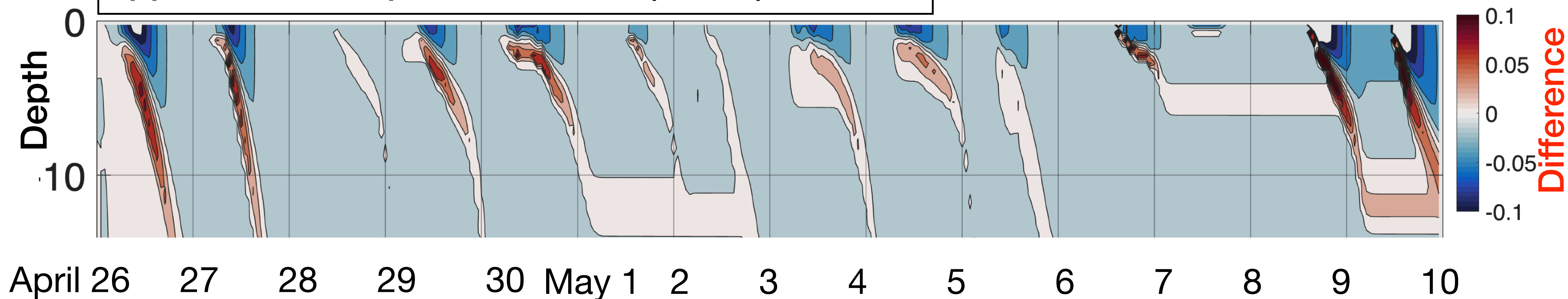
Upper ocean temperature response at AD08

using k- ϵ model

SST Difference : $T(\text{WBF}\theta) - T(\text{WBF})$

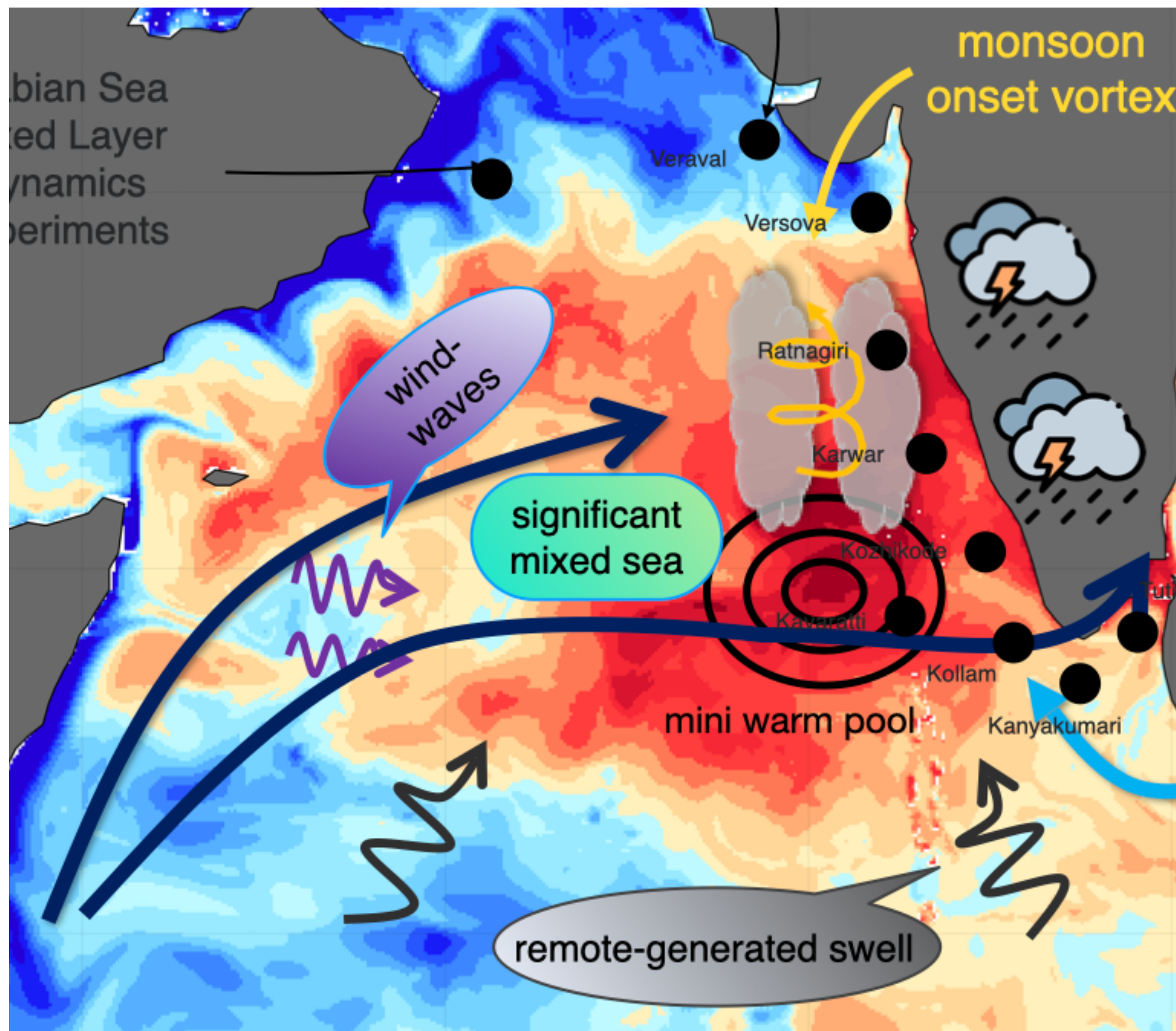


Upper ocean temp. Difference : $T(\text{WBF}\theta) - T$



Summary:

Misaligned waves, air-sea fluxes, and mini-warm pool SST



- A ubiquitous feature (**Raj et al. 2025**)
- Representing the effects on surface drag and upper-ocean turbulence challenging.
 - Especially in regions of strong density gradients and stratification
- Misaligned waves increase surface drag
 - Leading to enhanced turbulent mixing on diurnal time scales.
 - No publically available dataset yet; impossible to validate this.
 - Rectified effects on onset vortex
- More DCF data are needed.
 - Jim Edson: “COARE is not designed for low latitudes.”